ELSEVIER

Contents lists available at ScienceDirect

Chemosphere

journal homepage: www.elsevier.com/locate/chemosphere



Use of lead isotopes to identify sources of metal and metalloid contaminants in atmospheric aerosol from mining operations



Omar I. Félix ^a, Janae Csavina ^a, Jason Field ^b, Kyle P. Rine ^a, A. Eduardo Sáez ^{a,*}, Eric A. Betterton ^b

HIGHLIGHTS

- Lead isotopes used for source apportionment of lead contamination at mining sites.
- Atmospheric particulate and soils exhibit lead and arsenic contamination.
- Contaminant soil deposition patterns can be determined from lead isotopic analysis.

ARTICLE INFO

Article history:
Received 10 June 2014
Received in revised form 7 November 2014
Accepted 18 November 2014
Available online 12 December 2014

Handling Editor: X. Cao

Keywords:
Lead isotopes
Dust and aerosol
MOUDI
Smelting
Metal and metalloid contamination

ABSTRACT

Mining operations are a potential source of metal and metalloid contamination by atmospheric particulate generated from smelting activities, as well as from erosion of mine tailings. In this work, we show how lead isotopes can be used for source apportionment of metal and metalloid contaminants from the site of an active copper mine. Analysis of atmospheric aerosol shows two distinct isotopic signatures: one prevalent in fine particles (1 μ m aerodynamic diameter) while the other corresponds to coarse particles as well as particles in all size ranges from a nearby urban environment. The lead isotopic ratios found in the fine particles are equal to those of the mine that provides the ore to the smelter. Topsoil samples at the mining site show concentrations of Pb and As decreasing with distance from the smelter. Isotopic ratios for the sample closest to the smelter (650 m) and from topsoil at all sample locations, extending to more than 1 km from the smelter, were similar to those found in fine particles in atmospheric dust. The results validate the use of lead isotope signatures for source apportionment of metal and metalloid contaminants transported by atmospheric particulate.

© 2014 Elsevier Ltd. All rights reserved.

1. Introduction

Metal- and metalloid-laden dust and aerosol particles are produced during mining operations in a wide range of sizes. Coarse particles (>1 μ m aerodynamic diameter) are produced mainly by mechanical action like grinding and wind erosion (Csavina et al., 2011) while fine particles (<1 μ m) are the result of molten ore processing and smelting, being produced by condensation and coagulation of vapors (Jacob, 1999; Banic et al., 2006). The impact of airborne particles on human health depends on particle diameter and composition. Particles larger than 10 μ m are generally filtered out by the upper respiratory tract while particles smaller than 10 μ m (PM₁₀) are inhaled and may be transported to the lungs where they are phagocytized by alveolar macrophages and transported to the blood stream (Zheng et al., 2004). The

composition of the particles also plays a key role in health effects. Of particular concern is the presence of lead and other contaminants (e.g., arsenic) in atmospheric dust and aerosol that result from mining activities (Csavina et al., 2011, 2012; Mackay et al., 2013).

The atmosphere is the major initial recipient of lead among all environmental compartments. Global anthropogenic sources of lead are at least one order of magnitude higher than natural sources (Komárek et al., 2008). The determination of potential sources of atmospheric lead is important since its production can be controlled in advance, minimizing unwanted exposures to this toxic metal. However, it is difficult to discriminate between sources of lead and other metal and metalloid contaminants by only knowing their total concentration (Hopper et al., 1991). For this reason, lead isotopic analysis has been introduced as a "finger-printing" technique for lead contamination. Each source of lead may have a specific isotopic composition and the differences in this composition may be used to discriminate between possible sources (Komárek et al., 2008).

^a Department of Chemical and Environmental Engineering, University of Arizona, Tucson, AZ, USA

^b Department of Atmospheric Sciences, University of Arizona, Tucson, AZ, USA

^{*} Corresponding author. Tel.: +1 (520) 6215369. E-mail address: esaez@email.arizona.edu (A.E. Sáez).

The use of lead isotopes to determine possible sources has been widely reported (e.g., Grousset et al., 1994; Munksgaard and Parry, 1998; Veysseyre et al., 2001; Kurkjian et al., 2002; Zheng et al., 2004; Chen et al., 2005). Lead has four stable isotopes (average molar abundances in parenthesis): ²⁰⁴Pb (1.4%), ²⁰⁶Pb (24.1%), ²⁰⁷Pb (22.1%) and ²⁰⁸Pb (52.4%). ²⁰⁴Pb is the only non-radiogenic isotope and its abundance has remained constant over time (Russell and Farquhar, 1960). The abundances of ²⁰⁶Pb, ²⁰⁷Pb and ²⁰⁸Pb in minerals vary because they are derived from the decay of ²³⁸U, ²³⁵U and ²³²Th, respectively (Adgate et al., 1998). Isotopic composition of the radiogenic isotopes varies in different ores, depending on the age of their geological formation (Faure, 1986). An advantage of the use of lead isotopes for source apportionment is the fact that isotopic fractionation does not happen in industrial or environmental processes (Ault et al., 1970), owing mainly to the high atomic mass of lead. Lead isotopic composition is often expressed in terms of the concentration ratios 208Pb/206Pb and ²⁰⁷Pb/²⁰⁶Pb (Zheng et al., 2004; Chen et al., 2008).

Cumming and Richards (1975) observed that a plot of the concentration ratios ²⁰⁸Pb/²⁰⁶Pb vs. ²⁰⁷Pb/²⁰⁶Pb led to approximately a single monotonically increasing curve for all mineral samples analyzed. The radioactive decay of ²³⁸U, ²³⁵U and ²³²Th in terms of their relative abundances and half-lives has led to an increase in the concentration of ²⁰⁶Pb over time. As a consequence, the ratios ²⁰⁷Pb/²⁰⁶Pb and ²⁰⁸Pb/²⁰⁶Pb have decreased over time (Mukai et al., 2001). Fig. 1 shows the plot devised by Cumming and Richards, often called "growth curve", and the distribution along the curve of three different ores: The Iron King mine (Arizona, USA) is the oldest ore of those shown with an age of 1800 million years (Anderson and Guilbert, 1979), the Mount Isa ore (Queensland, Australia) has an age of approximately 1600 million years (Cumming and Richards, 1975) and the Ray Mine ore (Arizona, USA) has an age of 70 million years (Bouse et al., 1999).

In this work, we sampled atmospheric dust and aerosol at two different sites in Arizona: The first site is located in the towns of Hayden and Winkelman, where a copper mine and smelter currently operate; while the second site is located in Tucson, representing a local urban environment. The main objectives of this work were:

 To assess if lead isotopic analysis can be used for source apportionment of airborne particular matter associated with mining operations.

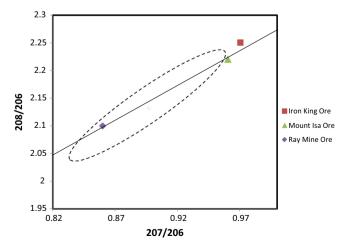


Fig. 1. The Cumming and Richards growth curve (solid line) for lead isotope concentration ratios (adapted from Chen et al., 2008), showing three mineral ores. The dashed region includes data from extensive sources (86 different sites in 35 countries) reported by Bollhöfer and Rosman (2001).

2. To investigate the extent of soil contamination in the vicinity of the smelting complex.

2. Materials and methods

2.1. Study site

The Hayden site is located approximately 80 km northeast of Tucson, Arizona. It is comprised of two towns: Hayden and Winkelman, with a combined population of approximately 1200. Currently, the site includes a concentrator, a copper smelter and tailings facilities. It is located at the confluence of the Gila and San Pedro Rivers. The processed copper ore originates from various mines in the vicinity, including the Ray mine. The Ray mine consist mainly of a porphyry copper deposit hosted in Pinal schist and Pioneer shale (w.mindat.org/loc-3377.html). In 2005, soil analysis showed that arsenic, lead and copper levels exceeded their respective residential soil remediation levels (EPA, 2012). The Environmental protection agency has reported elevated concentrations of arsenic, lead, copper, chromium and cadmium in atmospheric air samples in Hayden and Winkelman (EPA, 2012).

2.2. Sampling

The sampling site was located on the roof of a single-story High School building, approximately 2 km E of the mine tailings impoundment and 1 km SSE of the smelter (Fig. 2). A weather station and data logger (CR800, Campbell Scientific) providing information about temperature, relative humidity, wind speed and wind direction, was also present at the site. Wind speed and wind direction were analyzed with WindRose Pro (Enviroware) software. In Tucson, AZ, the sampling site was located on the roof of the five-story Physics and Atmospheric Sciences building at the University of Arizona.

A ten-stage micro-orifice uniform deposit impactor (MOUDI: M110-R, MSP Corporation) (Marple et al., 1991) was used to collect atmospheric dust and aerosol. The MOUDI was operated at a flow rate of 30 L min⁻¹ for 96-h sampling periods. Sampling was performed in two modes: regular sampling was done by operating the MOUDI continuously for 96 h, while programmed sampling was performed by operating the MOUDI only for specified wind speeds and directions until total operating time reached 96 h. The calibrated cut points (d_{50} -values) for the MOUDI sampler are 18, 9.9, 6.2, 3.1, 1.8, 1.0, 0.55, 0.32, 0.18, 0.10 and 0.054 μm equivalent aerodynamic diameters. Teflon substrates (PTFE membrane, 2-µm pore size, 46.2 mm diameter, Whatman) were used for sample collection. The substrates were cleaned sequentially by DI water and methanol before use. They were transported to and from the field site in enclosed impactor holders (MSP Corporation). Substrates were weighed before and after sampling using EPA class I equivalent methods on an ultra-microbalance (Mettler Toledo XP2U). Aluminum substrates (47 mm, MSP Corporation) were occasionally used for electron microscope analysis samples.

Soil samples were taken at Hayden at four different distances from the smelting complex in a straight line transect NE of the smelter, a prevailing wind direction (Fig. 3). Samples were collected at different depths to obtain a vertical profile of the contamination: 0–3 mm, 3–6 mm, 6–9 mm, at 50 mm, and finally at 100 mm below the surface. Final samples at each site and depth were obtained by mixing three different samples. The composite samples were dried for 10 h at a temperature of 110 °C and then sieved (ASTM, 2010) through a 0.84 mm sieve to discard coarser fractions.



Fig. 2. Satellite picture of the sampling site in Hayden, AZ, with locations of mining and sampling operations. Source: Google Earth.



Fig. 3. Satellite picture showing the four soil sampling locations (labeled A–D) at the Hayden site. The shadow of the 305 m high smokestack can be seen towards the WNW. Approximate distances of the sampling points from the smelter location: A – 650 m, B – 900 m, C – 1300 m, and D – 4000 m. Altitudes with respect to sea level of sampling points are: A – 680 m, B – 760 m, C – 710 m, and D – 620 m. Source: Google Earth.

2.3. Sample preparation

Exposed Teflon substrates were transferred to glass vials and extracted with 15 mL of aqua regia (1.03 M HNO $_3$ /2.23 M HCl, trace-metal grade) in a sonicator at 80 °C for 60 min (Harper et al., 1983). Aliquots of 1.2 mL were diluted to 4 mL with deionized water before analysis. Soil samples were extracted in the same way, using 5 mg of soil in each sample. Prior to lead isotope analysis, samples were concentrated on a hot plate. The extraction method was verified by comparing it with the use of concentrated HNO $_3$, or a combination of concentrated HNO $_3$ /H $_2$ O $_2$ solutions in conjunction with microwave digestion, all of which yielded similar concentrations for representative samples (Gonzales et al., 2014), which suggests that element measures represent total acid extractable concentrations.

2.4. Sample analysis

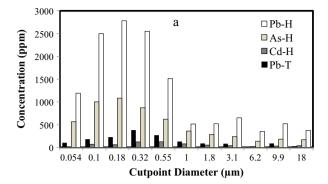
An ICP-MS (Agilent 7700X with Octopole reaction system) was used for the determination of lead, arsenic and cadmium, as well as lead isotopes. The equipment was tuned for robust plasma conditions to reduce the formation of oxides to less than 2%. Certified calibration standards from Accustandard were prepared with MiliQ

water, 0.669 M HCl (Fisher, trace-metal grade) and 0.309 M HNO₃ (EMD, Omnitrace). National Institute of Standards and Technology (NIST) standard reference material (SRM 1643e trace elements in water) was analyzed with each set of samples. NIST SRM 981 (lead isotopic standard) was used for calibration of lead isotope measurements. The analytical precision of lead isotopic ratios was under 0.5% for the concentration ratios Pb²⁰⁷/Pb²⁰⁶ and Pb²⁰⁸/Pb²⁰⁶.

Morphologic and elemental analyses were done at University Spectroscopy and Imaging Facilities of the University of Arizona using a field emission scanning electron microscope (Hitachi S-4800 Type II SEM) coupled to energy dispersive spectroscopy (ThermoNORAN NSS EDS). Aluminum substrates were used in the MOUDI when collecting samples for this purpose to improve resolution.

3. Results and discussion

Fig. 4a shows a comparison of concentrations in atmospheric particulate at the two field sites for a representative sampling period. Contaminant concentrations follow a bimodal distribution at the Hayden site with maxima around 0.18 and 9.9 μ m. This trend was consistently seen over sampling periods spanning several



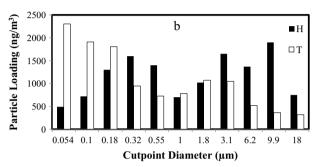


Fig. 4. (a) Concentrations (wt/wt) of As, Cd and Pb as a function of particle size from representative (consecutive 96-h operation) MOUDI samples of dust and aerosol in atmospheric air taken at the two sites (H – Hayden site, and T – Tucson site). Arsenic and cadmium were below detection limits in the Tucson samples. (b) Mass concentration of total particles as a function of particle size for the sample in (a).

years of sampling (Csavina et al., 2011). The highest concentration of contaminants is present in the accumulation mode (particles between 0.1 and 1 μm) (Seinfeld and Pandis, 2006). Smelting and high-temperature processes release submicron particles formed by condensation of vapors, which coalesce to enrich the accumulation mode. Note that both contaminant concentration and particle mass concentration follow a bimodal distribution at the Hayden site.

Lead and arsenic concentrations of the mine tailings impoundment in Hayden are 14.0 ± 6.8 and 46.4 ± 33.7 ppm (EPA, 2012), respectively, which suggests that the main source of atmospheric contamination at this site is smelting activity and not the erosion of mine tailings, given the high contaminant concentrations in atmospheric particulate (100-1000 ppm, Fig. 4a).

Lead isotope analysis was performed at the two sites of study. Particles in the range of 0.32–0.55 um (denoted as fine particles here) and 3.1-6.2 um (coarse particles) were analyzed to obtain the lead isotopic composition. Fig. 5 shows the lead isotope ratios associated with two sampling periods at Hayden, along with the corresponding wind roses. Two different isotopic signatures for coarse and fine particles are evident (Fig. 5a), which suggests the existence of two different atmospheric lead sources. The isotopic ratios of the fine particles coincide with those of the Ray Mine (Fig. 1) which is the main ore source for the smelter. This is evidence that the lead found in the fine particles is associated with condensation of high-temperature vapors produced at the smelting site. The different isotopic signature of the coarse particles suggests a different origin, possibly related to the background lead present in the area (see below). It is interesting to note that the two sampling periods shown in Fig. 5 corresponded to completely

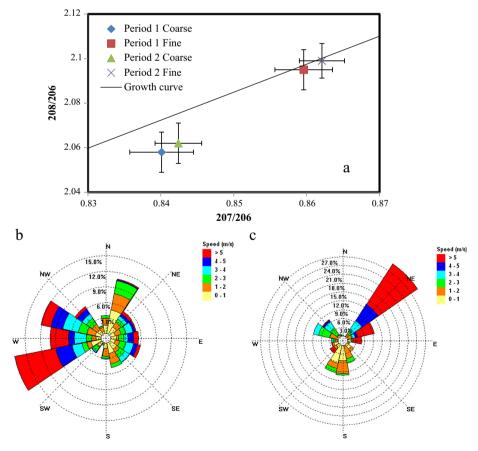


Fig. 5. (a) Lead isotope concentration ratios obtained during MOUDI sampling for two 96-h sampling periods at the Hayden site. Error bars show standard deviations of triplicate measurements by ICP-MS; (b) wind rose for period 1; (c) wind rose for period 2.

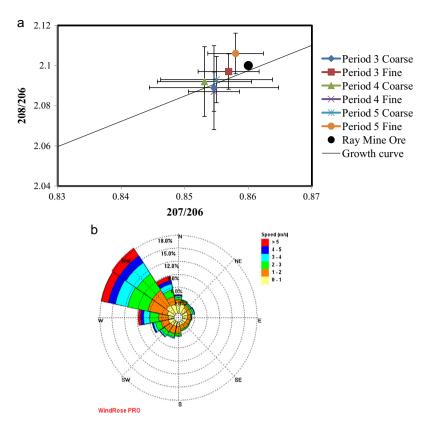


Fig. 6. (a) Lead isotopic ratios obtained from MOUDI samples during programmed sampling when wind direction was from the smelter to the sampling site (300–360°) at the Hayden site; (b) representative wind rose of the programmed sampling periods.

different average wind directions. In fact, the concentration of lead in the fine particles during period 1 (12.5 ng/m³, wind from WSW) was around twice the concentration for period 2 (6.2 ng/m³, wind from NE). Note that during period 1, wind patterns favor the direction from the smelter to the sampling site.

Three different programmed sets of samples were taken with the MOUDI at the Hayden site (periods 3-5). The MOUDI was programmed to sample only when the wind was coming from the smelting area (300-360°). Results obtained from these samples, as well as a representative wind rose of the sampling periods are presented in Fig. 6. Fig. 6a shows that both coarse and fine particles exhibit only one isotopic signature with high ratios similar to the results shown on Fig. 5a for fine particles. This indicates that for the programmed samples the smelter was the predominant source. Note that the isotopic ratios for all the sampling periods are within experimental error, which suggests that smelter emissions have a consistent signature, that is, the ore processed in the smelter seems to have the same origin or, at least, the same lead isotopic composition. This does not mean dust and aerosol from other sources were absent. However, since the lead concentration is appreciably higher for samples coming from the smelter area, their lead isotopic signature predominates. We also note that due to the complex topography of the Hayden site, there is no certainty that winds followed the general line of sight (300–360°).

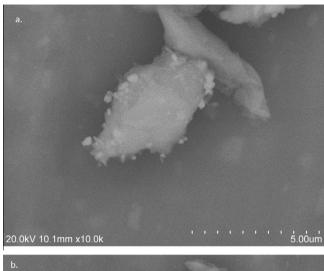
The programmed samples were analyzed by scanning electron microscopy (SEM). Sample of images obtained are shown in Fig. 7. Deposition of fine particles onto coarse particles is clearly observed in Fig. 7a. Energy dispersive spectroscopy (EDS) was used to determine the content of lead in the fine particles attached; it was found that these particles contain lead in concentrations as high as 14.85 ± 1.88% by weight. Fig. 7b shows coarse particles in which the attachment of fine particles is absent. The lead content

of these particles was negligible, as determined by EDS. This confirms that the second lead signature seen in the MOUDI samples collected without discriminating for wind direction (coarse particles in Fig. 5a) is, in this case, masked by the prevalence of lead-enriched aerosol in both fine and coarse particles.

The coarse particle lead ratios (Fig. 5a) were similar to those found in samples taken at the Tucson site, which is free of mining activity (Fig. 8). Coarse particles may be associated with wind-blown soil from the site. For Tucson, coarse and fine particles had the same lead isotopic ratios. This result suggests that the Tucson lead, which coincides with the Hayden lead in coarse particles, is unaffected by smelting activities, and it might be representative of a background source of lead common to the region, presumably legacy lead from fuel sources, since this isotopic signature is similar to that of lead used as a gasoline additive, which originates from Missouri ores: ²⁰⁸Pb/²⁰⁶Pb = 2.07, ²⁰⁷Pb/²⁰⁶Pb = 0.84 (Sutherland et al., 2003; Bollhöfer and Rosman, 2001).

The results presented above point to a significant atmospheric contamination at the Hayden site that originates from smelting activities. We investigated how this contamination affects the soil in the vicinity of the smelter. Soil samples were analyzed for Pb, As and Cd at various distances from the smelter and various depths. Contaminant concentrations are shown in Figs. 9 and 10. As expected, contaminant concentrations in topsoil decreased monotonically with distance from the smelter (Fig. 10).

Lead isotopic analyses were performed on soil samples A–C (Fig. 3) to determine the source of contamination and possible effect of the deposition of smelter fine particles on the soil. Sample D was not analyzed for lead isotopes due to the low concentration of lead present. For each of the samples analyzed, only the top and the bottom layers were considered. Fig. 11 shows the isotopic ratios obtained for the different soil samples. Isotopic ratios



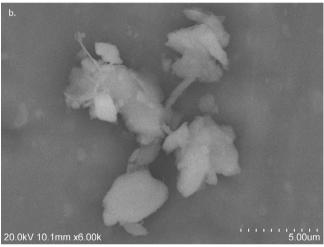


Fig. 7. Scanning electron microscope images from MOUDI samples from period 3: (a) some coarse particles show fine particles attached to them, while (b) other coarse particles are free of fine particles.

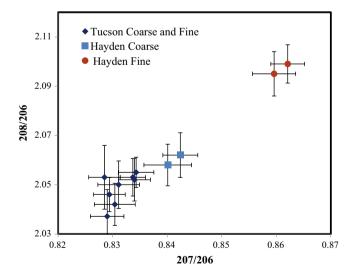


Fig. 8. Comparison between lead isotopic ratios obtained in atmospheric dust and aerosol in Hayden and Tucson, AZ.

obtained in the sample closest to the smelter (sample A) are similar to the ratios obtained for the airborne fine particles in the atmosphere at Hayden (Fig. 5). Similar ratios were observed in the 100 mm deep sample, which indicates a similar source of lead

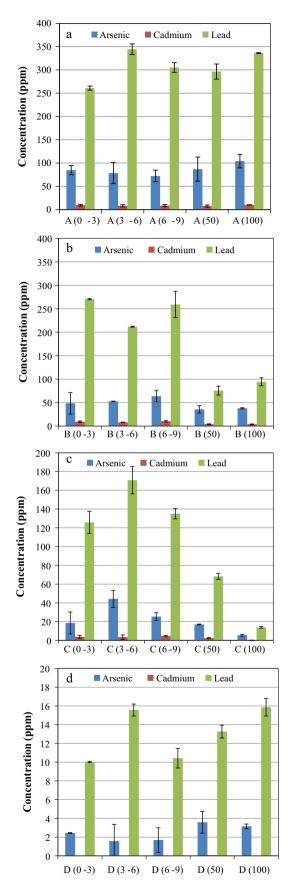


Fig. 9. Total lead, cadmium and arsenic in Hayden site soil samples; distance from the smelter: (a) 650 m; (b) 900 m; (c) 1300 m; (d) 4000 m. Horizontal axis values represents range of depth of samples in mm.

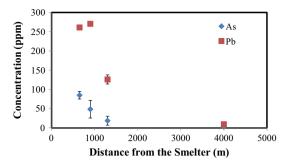


Fig. 10. Lead and arsenic concentrations in topsoil (0-3 mm depth) at the Hayden site as a function of distance from the smelter.

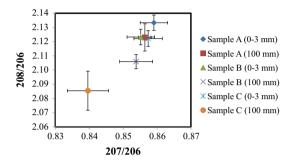


Fig. 11. Isotopic ratios for soil samples in Hayden. Distance from the smelter specified in Fig. 3. Numbers in parentheses represent the depth of the sample.

throughout the soil profile analyzed. Sampling point A is located only 650 m from the smelter. These results indicate that, despite weathering and rain runoff, lead and arsenic accumulate in the soil, and the contamination extends uniformly through at least 10 cm of depth (Fig. 9a). Lead isotope results provide strong evidence that the contamination originates in atmospheric transport of smelter emissions.

Isotopic ratios for the topsoil of sample B (Fig. 11) are again close to the smelter emissions. However, the ratios for the deepest sample (100 mm) definitely indicate the presence of a different source mixed with the emissions source. This trend becomes more evident for sample C (1300 m), for which the topsoil layer still exhibits the smelter isotopic ratios, but the 100 mm depth sample shows significantly lower ratios, which trend in the direction of the presumed background measured in the coarse atmospheric particles (compare with Fig. 5a).

4. Conclusion

The results obtained in this work demonstrate that lead isotope analysis can be used for source apportionment of metal and metalloid contaminants from mining operations. At the mining site in Hayden, AZ, two different lead sources were identified: one associated with fine particles originating from condensation of high temperature vapors emitted at the smelter site, and another one associated with a regional background, which shows the same isotopic composition as legacy lead used in fuels. Topsoil samples near the smelting site show concentrations of Pb and As decreasing with distance from the smelter but lead isotope ratios similar to those found in fine particles in atmospheric dust, which are characteristic of smelting emissions. On the other hand, isotope ratios at a soil depth of 100 mm decrease with distance from the smelter. These results are consistent with a contaminant deposition pattern from windblown particles from the smelter site to surrounding soils. These results validate the use of lead isotope signatures for source apportionment of metal and metalloid contaminants transported by atmospheric particulate.

Acknowledgements

This work was supported by Grant number P42 ES04940 from the National Institute of Environmental Health Sciences (NIEHS), and National Institutes of Health (NIH). The views of authors do not necessarily represent those of the NIEHS, NIH.

References

Adgate, J.L., Rhoads, G.G., Lioy, P.J., 1998. The use of isotope ratios to apportion sources of lead in Jersey City, NJ, house dust wipe samples. Sci. Total Environ.

Anderson, P., Guilbert, I.M., 1979, Precambrian massive sulfide deposits of Arizona – a distinct metallogenic epoch and province. Nev. Bur. Mines. Geol. Rep. 33, 39-48

ASTM (American Society for Testing and Materials). Standard test methods for laboratory determination of water (moisture) content of soil and rock by mass. Report D2216-10, 2010.

Ault, W.U., Senechal, R.G., Erlebach, W.E., 1970. Isotopic composition as a natural tracer of lead in the environment, Environ, Sci. Technol, 4, 305-313.

Banic, C.M., Leaitch, W.R., Strawbridge, K.B., Tanabe, R., Wong, H.K.T., Gariepy, C. Simonetti, A., Nejedly, Z., Campbell, J.L., Lu, J., 2006. The physical and chemical evolution of aerosols in smelter and power plant plumes: an airborne study. Geochem.: Explor. Environ. Anal. 6, 111-120.

Bollhöfer, A., Rosman, K.J.R., 2001. Isotopic source signatures for atmospheric lead: the northern hemisphere. Geochim. Cosmochim. Acta 65, 1727-1740.

Bouse, R.M., Ruiz, J., Titley, S.R., Tosdal, R.M., Wooden, J.L., 1999. Lead isotope compositions of late Cretaceous and early Tertiary igneous rocks and sulfide minerals in Arizona: implications for the sources of plutons and metals in porphyry copper deposits. Econ. Geol. 94, 211-244.

Chen, J., Tan, M., Li, Y., Zhang, Y., Lu, W., Tong, Y., Zhang, G., Li, Y., 2005. A lead isotope record of Shanghai atmospheric lead emissions in total suspended particles during the period of phasing out of leaded gasoline. Atmos. Environ. 39 1245-1253

Chen, J., Tan, M., Li, Y., Zheng, J., Zhang, Y., Shan, Z., Zhang, G., Li, Y., 2008. Characteristics of trace elements and lead isotope ratios in PM(2.5) from four sites in Shanghai. J. Hazard. Mater. 156, 1-3.

Csavina, J., Landazuri, A., Wonaschutz, A., Ryne, K., Rheinheimer, P., Barbaris, B., Conant, W., Sáez, A.E., Betterton, E.A., 2011. Metal and metalloid contaminants in atmospheric aerosols from mining operations. Water Air Soil Poll. 221, 145-

Csavina, J., Field, J., Taylor, M.P., Gao, S., Landazuri, A., Betterton, E.A., Sáez, A.E., 2012. A review on the importance of metals and metalloids in atmospheric dust and aerosol from mining operations. Sci. Total Environ. 433, 58-73.

Cumming, G.L., Richards, J.R., 1975. Ore lead isotope ratios in a continuously changing Earth. Earth Planet. Sci. Lett. 28, 155-171.

EPA (Environmental Protection Agency). Asarco Hayden Plant. http:// yosemite.epa.gov/r9/sfund/r9sfdocw.nsf 7508188dd3c99a2a8825742600743735/

3940634a9aec311e88257478006736ce!OpenDocument> (accessed 2012).

Faure, G., 1986. Principles of Isotope Geology. Wiley, New York.

Gonzales, P., Felix, O., Alexander, C., Lutz, E., Ela, W., Sáez, A.E., 2014. Laboratory dust generation and size-dependent characterization ofmetal and metalloidcontaminated mine tailings deposits. J. Hazard. Mater. 280, 619-626.

Grousset, F.E., Quetel, C.R., Thomas, B., Buat-Menard, P., Donard, O.F., Bucher, A., 1994. Transient Pb isotopic signatures in the western European atmosphere. Environ. Sci. Technol. 28, 1605-1608.

Harper, S.L., Walling, J.F., Holland, D.M., Pranger, L.J., 1983. Simplex optimization of multielement ultrasonic extraction of atmospheric particulates. Anal. Chem. 55,

Hopper, J.F., Ross, H.B., Sturges, W.T., Barrie, L.A., 1991. Regional source discrimination of atmospheric aerosols in Europe using the isotopic composition of lead. Tellus 43, 45-60.

Jacob, D.J., 1999. Introduction to Atmospheric Chemistry. Princeton University Press, Princeton.

Komárek, M., Ettler, V., Chrastný, V., Mihaljevič, M., 2008. Lead isotopes in environmental sciences: a review. Environ. Int. 34, 562-577.

Kurkjian, R., Dunlap, C., Flegal, A.R., 2002. Lead isotope tracking of atmospheric response to post-industrial conditions in Yerevan, Armenia. Atmos. Environ. 36,

Mackay, A.K., Taylor, M.P., Munksgaard, N.C., Hudson-Edwards, K.A., Burn-Nunes, L., 2013. Identification of environmental lead sources and pathways in a mining and smelting town: Mount Isa, Australia. Environ. Poll. 180, 304-311.

Marple, V.A., Rubow, K.L., Behm, S.M., 1991. A microorifice uniform deposit impactor (MOUDI): description, calibration, and use. Aerosol Sci. Technol. 14,

Mukai, H., Machida, T., Tanaka, A., Vera, Y.P., Uematsu, M., 2001. Lead isotope ratios in the urban air of eastern and central Russia. Atmos. Environ. 35, 2783-2793.

- Munksgaard, N.C., Parry, D.L., 1998. Lead isotope ratios determined by ICP-MS: monitoring of mining-derived metal particulates in atmospheric fallout, Northern Territory, Australia. Sci. Total Environ. 217, 113–125.
- Russell, R.D., Farquhar, R.M., 1960. Lead Isotopes in Geology. Interscience Publishers, New York.
- Seinfeld, J.H., Pandis, S.N., 2006. Atmospheric Chemistry and Physics. From Air Pollution to Climate Change. Wiley, New York.
- Sutherland, R.A., Day, J.P., Bussen, J.O., 2003. Lead concentrations, isotope ratios, and source apportionment in road deposited sediments, Honolulu, Oahu, Hawaii. Water Air Soil Poll. 142, 165–186.
- Veysseyre, A.M., Bollhöfer, A.F., Rosman, K.J., Ferrari, C.P., Boutron, C.F., 2001.
 Tracing the origin of pollution in French alpine snow and aerosols using lead isotopic ratios. Environ. Sci. Technol. 35, 4463–4469.
 Zheng, J., Tan, M., Shibata, Y., Tanaka, A., Li, Y., Zhang, G., Zhang, Y., Shan, Z., 2004.
- Zheng, J., Tan, M., Shibata, Y., Tanaka, A., Li, Y., Zhang, G., Zhang, Y., Shan, Z., 2004. Characteristics of lead isotope ratios and elemental concentrations in PM10 fraction of airborne particulate matter in Shanghai after the phase-out of leaded gasoline. Atmos. Environ. 38, 1191–1200.